Real World Globes – Exploring Deep-Ocean Water Masses

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Purpose:

- To discuss the global "thermohaline circulation", the mechanisms by which it operates, and the measures implemented to quantify water mass age.
- To perform basic plotting on a gridded surface, and to recognize patterns inherent within the plotted data.
- To practice basic conversions and calculations when provided a formula and values.

Target Audience:

- Non-science Major Undergraduates
- Science Major Undergraduates

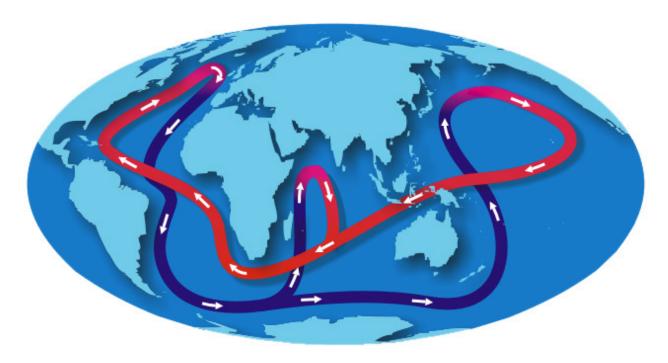
Materials:

- Mother Earth GlobeTM
- Great Circle ruler and grid
- Transparent Atlantic Salinity Overlay
- Dry-erase markers, eraser, and calculator

Introduction:

The world's oceans are layered, separated by differences in temperature and salinity. This stratification results in a differentiation of large bodies of the deep ocean into separate *water masses*, specifically defined by their temperature and salinity values. These water masses acquire their distinctive characteristics from their point of origin: either from the North Atlantic within the Norwegian/Greenland Seas, or from the Weddell Sea off the coast of Antarctica. At these two regions of the globe, natural processes increase the density of the ocean water that sits at the surface. In consequence, the surface water sinks into the deep ocean, where it begins its journey around the world's oceans as a part of the interconnected global "conveyor belt" of ocean currents known as the *thermohaline circulation* (see picture below).

As well as temperature and salinity, defining chemical characteristics, such as radiocarbon (¹⁴C) concentrations, create noticeable distinctions between different water masses. ¹⁴C, a radioactive variant of carbon with a half-life approximately 5,700 years, is used to *carbon date* the age of these water masses, which tells us how long it takes to make one trip around the "conveyor belt". ¹⁴C is produced in the atmosphere, transformed from elemental nitrogen by cosmic radiation. When surface waters are in contact with the atmosphere, the amount of ¹⁴C in the ocean equilibrates with the atmosphere. Once those surface waters sink into the deep ocean, that water mass stops acquiring new ¹⁴C, and the ¹⁴C already in the water mass begins to decay. This is analogous to starting a stopwatch the minute the water ceases contact with the atmosphere; from that point onward, the time since the water was at the surface of the ocean can be calculated. This is what we refer to when we speak of the water's "age".

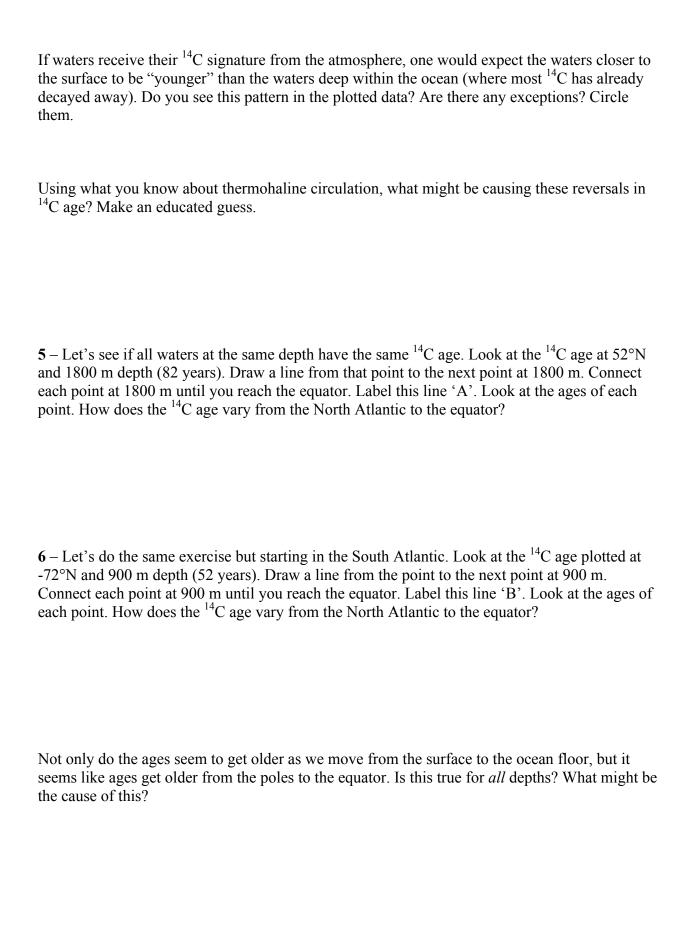


The global thermohaline circulation begins with the sinking of cold, salty water in the North Atlantic. From there, the water travels south, around Antarctica, and eventually into the Indian and North Pacific Oceans, where it rises to the surface and begins the journey back to the Atlantic.

Activity:

An oceanographic vessel has just returned from two months at sea. The research team traveled down the center of the Atlantic Ocean from Rekjavík, Iceland to South Georgia Island near the Antarctic Peninsula. Along their transect of the Atlantic Ocean, they stopped at a number of stations and took vertical measurements of the ¹⁴C concentration.

- 1 Assemble the globe. Place the grid atop the Great Circle and orient the circle so that it travels through the Atlantic Ocean.
- **2** Label the x-axis of the Great Circle grid "Latitude (°)" and label each vertical line with latitude values. (The center of the grid should be labeled 0° due to its position on the equator one box on the grid is equivalent to 2°). **Remember:** Positive latitude values are to the north of the equator, where negative latitude values are to the south. Also, label the y-axis "Depth (m)" and label each horizontal line from the outside of the circle inward (beginning with 0 at the surface each box is equal to 300 m).
- 3 Orient the Great Circle with superimposed grid so that it aligns down the center of the Atlantic Ocean.
- **4** Plot the points listed in Table 1 on the grid and label the ¹⁴C age next to each point.



Additional Activity for Science Major Undergraduates:

Knowing the age of two parcels of water at two given locations allows us to calculate the velocity at which the water is flowing. Knowing how fast water masses are flowing tells us something about how fast the larger thermohaline circulation is moving (which is important for calculating ocean mixing times, heat transport, and a bunch of other pieces of important oceanographic information that we're not focused on here). We will be calculating the velocity of **NADW** along line 'A' and Antarctic Intermediate Water (**AAIW**) along line 'B'. **Remember:** Velocity is *distance* divided by *time*, often reported as meters per second (m/s).

Make sure to show ALL of your work on the attached page!!!

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8 – First, we need to convert our latitude coordinates in <u>degrees</u> to distance measurements in <u>meters</u> . Find the distance in degrees between the first and last points of line 'A'. Using the conversion that $1^{\circ} = 111,000$ m, how far apart are the two points?
9 – Now, we must compute the difference in age between the last point and the first point and convert that into seconds. Using the conversion that 1 year = $31,536,000$ s, what is the time difference between the two points?
10 – Finally, take the answer from question 8 and divide it by the answer from question 9. What is the velocity of NADW ?
11 – Now let's do the same steps for line 'B'. Compute the distance between the first and last points, calculate the time difference, and divide them just as outlined above. What is the velocity of AAIW ?
Is NADW or AAIW faster? Circle the correct answer.

SHOW ALL OF YOUR CALCULATIONS! DO NOT FORGET TO INCLUDE UNITS!

Table 1 – Radiocarbon Ages from Atlantic Ocean cruise

Latitude (°N)	Depth (m)	¹⁴ C Age (Years)	Latitude (°N)	Depth (m)	¹⁴ C Age (Years)
52	0	16	8	3600	478
52	900	57	-11	0	7
52	1800	82	-11	900	253
52	2700	98	-11	1800	573
40	0	23	-11	2700	599
40	900	92	-11	3600	362
40	1800	121	-26	0	27
40	2700	141	-26	900	196
40	3600	168	-26	1800	622
34	0	21	-26	2700	234
34	900	114	-26	3600	301
34	1800	207	-46	0	9
34	2700	225	-46	900	122
34	3600	229	-46	1800	715
34	4500	316	-46	2700	284
24	0	42	-46	3600	392
24	900	215	-60	0	16
24	1800	398	-60	900	69
24	2700 3600	401	-60	1800	111
24 8	3600 0	418 11	-60	2700	197
8	900	451	-60	3600	219
8	1800	462	-72	0	10
8	2700	402 471	-72	900	52